

**A REVIEW OF THE
METAMORPHIC
PETROLOGY OF THE
APPALACHIAN MOUNTAINS**

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2 Introduction:

3 The Appalachian mountains run along the entirety of the North American east coast (see Figure 1)
4 (Murphy, n.d.). The most northern part of the mountain belt is the oldest, with age decreasing to
5 the southwest (Murphy, n.d.). It also exhibits strong variation in metamorphic petrology
6 throughout the range due to it being made from multiple mountain-building events (USGS:
7 Geology of the Southern Appalachian Mountains, n.d.; Marshal University, 2010). Three main
8 orogenies occurred in order to make this polyorogenic belt: the Taconic orogeny, Acadian orogeny,
9 and Alleghanian orogeny (“USGS Geology and Geophysics,” n.d.; USGS: Geology of the
10 Southern Appalachian Mountains, n.d.; Marshal University, 2010; Rast, 1989) These orogenies
11 represent the collision of island arcs, the destruction of Paleozoic oceans, and the collision of

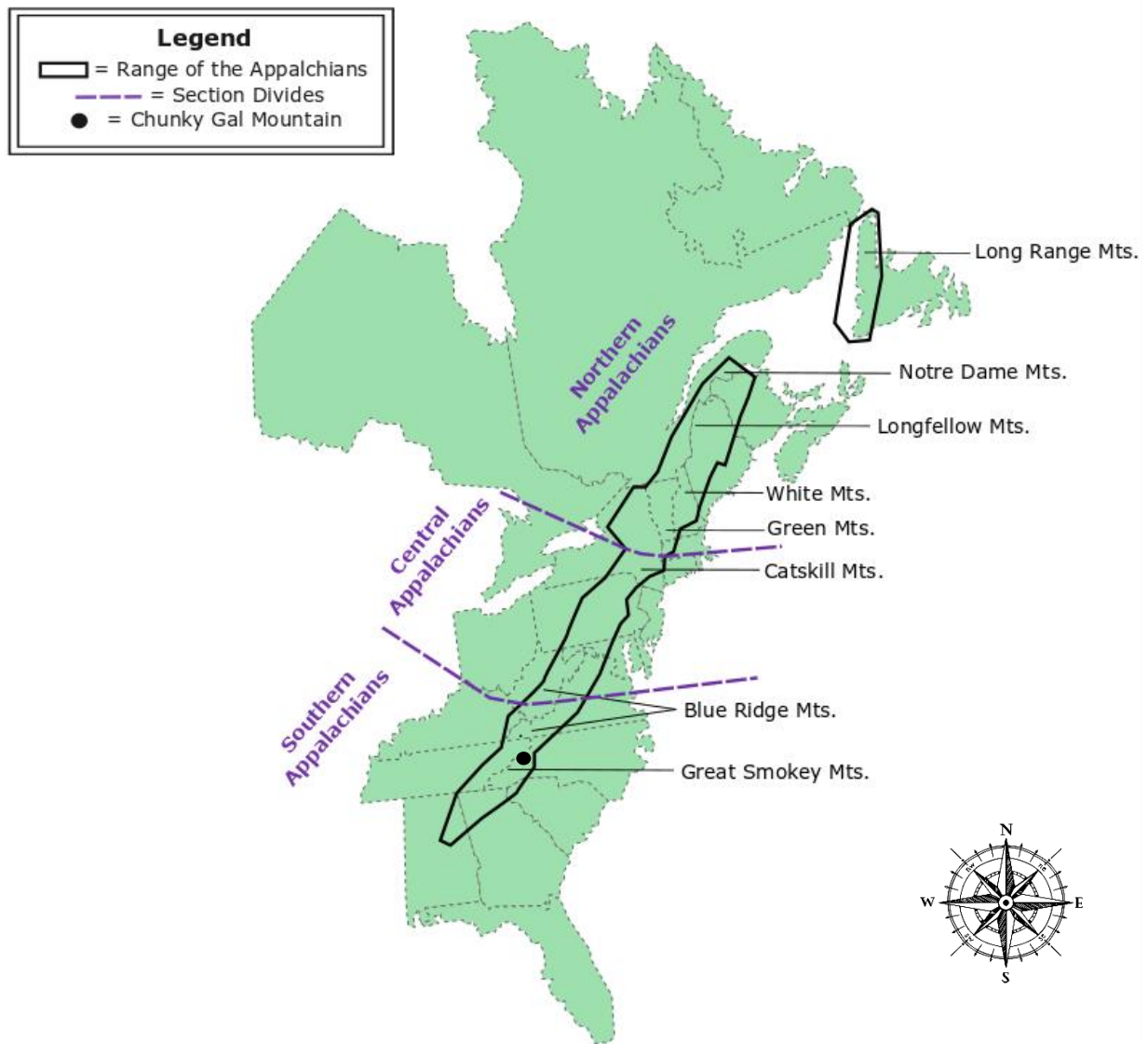


Figure 1: A map of North America showing the location of mountain ranges and sectional divides of the Appalachian Mountains.

A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

12 Africa with North America during the formation of Pangaea (USGS: Geology of the Southern
13 Appalachian Mountains, n.d.; Marshal University, 2010; Hatcher, 2002). The basement of these
14 orogenies is comprised of plutonic and metasedimentary rocks that metamorphosed into gneisses
15 (Rast, 1989). The majority of the northern Appalachians contains charnockites and paragneisses,
16 with the southern Appalachians being mainly orthogneisses (Rast, 1989).

17 **Northern Appalachians and the Taconic Orogeny:**

18 The Taconic orogeny formed the northern Appalachians in the Ordovician and Silurian (Murphy,
19 n.d.; Marshal University, 2010; Rast, 1989). It involved an extensive amount of thrusting and the
20 convergence of volcanic island arcs (Murphy, n.d.; Marshal University, 2010; Rast, 1989; Hatcher,
21 2002). The area that is most widely considered Taconian by researchers is the western margin of
22 the northern Appalachians, likely because it is the area least impacted by the other orogenic events
23 (Rast, 1989). As designated by Rast (1989), the Northern Appalachians are comprised of the Long
24 Range Mountains (Newfoundland), the Notre Dame Range (Quebec), the Longfellow Mountains
25 (Maine), the White Mountains (New Hampshire), and the Green Mountains (Vermont) (See Figure
26 1) (Slattery, Cobban, Mckinney, Harries & Sandness, 2013). In the northern Appalachians, the age
27 of the deformation decreases to the north, placing the youngest deformation in the Long Range
28 Mountains (Murphy, n.d.). It can be seen that higher temperatures and pressures quickly follow
29 the regional metamorphism of the Taconic orogeny in this section of the Appalachians (Dorfler,
30 Tracy, & Caddick, 2014).

31 The Catskill Mountains of New York, USA are the most publicized of the ranges in the
32 Northern Appalachians and show high-grade metamorphism succeeded by rapid uplift (Murphy,
33 n.d.; Dorfler, Tracy, & Caddick, 2014). Dorfler, Tracy, & Caddick (2014) described the contact
34 metamorphism in New York as pelitic schists which underwent the regional metamorphism of the
35 Taconian orogeny. This regional metamorphism was immediately followed by contact
36 metamorphism produced by mafic intrusions from the same orogenic event (Dorfler, Tracy, &
37 Caddick, 2014).

38 The first evidence of metamorphism in the samples is the staurolite-garnet-plagioclase Taconic
39 regional metamorphism. Calcium enrichment in plagioclase and garnet crystals is representative
40 of an intermittent cooling period between the stages of high-pressure crystal growth (Dorfler,
41 Tracy, & Caddick, 2014). While the regional metamorphism produced porphyritic and schistose
42 textures, contact metamorphism generated reaction textures. Samples that had undergone both
43 forms of metamorphism also possessed overprinting textures. In these textures, marks of earlier
44 processes were still visible but overlain by the new contact metamorphism (Dorfler, Tracy, &
45 Caddick, 2014). Evidence of overprinting in samples from the northern Appalachians included:

- 46 1. Garnet porphyroblasts with chemically modified rims (see Figure 2),
- 47 2. A difference in composition between the matrix and inclusion plagioclase/biotite,
- 48 3. And the presence of spinel within the matrix.

49 Samples that experienced both types of metamorphism contained quartz, biotite, plagioclase,
50 garnet, sillimanite, spinel, ilmenite, and graphite and lacked muscovite and alkali feldspar. Biotite,
51 plagioclase, quartz, and sillimanite are abundant inclusions within the garnet porphyroblasts.
52 Some minor inclusions of ilmenite can also be seen. Garnet, plagioclase, and biotite are often the
53 minerals used in thermobarometric calculations, as they best record the changes in metamorphic
54 conditions (Williams, 2011; Dorfler, Tracy, & Caddick, 2014).

A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

55 A large fluctuation in heat generated high enough temperatures to produce melting, allowing
56 the overprinting thermal metamorphism to remove the matrix muscovite and staurolite. This left
57 only the inclusion muscovite and staurolite and explains the depletion of these minerals within the
58 matrix. The gabbroic melts formed by the high temperatures also depleted the melt of water and
59 felsic components through the recrystallization of plagioclase and removal of muscovite (Dorfler,
60 Tracy, & Caddick, 2014). The heating initiated the breakdown of available micas, staurolite, and
61 garnet, releasing the necessary elements to produce sillimanite. This mineral assemblage could
62 only be produced by the heat from intruding magmas (Dorfler, Tracy, & Caddick, 2014). Since
63 there was no exchange of iron or magnesium between garnet and biotite crystals, it can be inferred
64 that there was rapid cooling from the high temperatures made by the gabbroic melts (Dorfler,
65 Tracy, & Caddick, 2014). In comparison to a collision of two continental edges, the mafic
66 intrusions and hydrous minerals within the samples are further indicative of a collision of island
67 arcs due to the amount of water present during formation (Dorfler, Tracy, & Caddick, 2014).

68 Garnet poikiloblasts in the northern Appalachians showed distinct zoning patterns that
69 indicated multiple thermal events. There are biotite, sillimanite, quartz, and ilmenite inclusions
70 in the inner crystal; however, the outer rim only has the rare inclusion of spinel. Electron probe
71 analysis by Dorfler, Tracy, & Caddick (2014) revealed the outermost rim of garnet poikiloblasts
72 to be enriched with calcium and the opposite to be true with magnesium, as seen in Figure 2. This
73 chemically modified rim is the result of a change the composition of surrounding material and
74 minerals produced, indicating high temperatures.

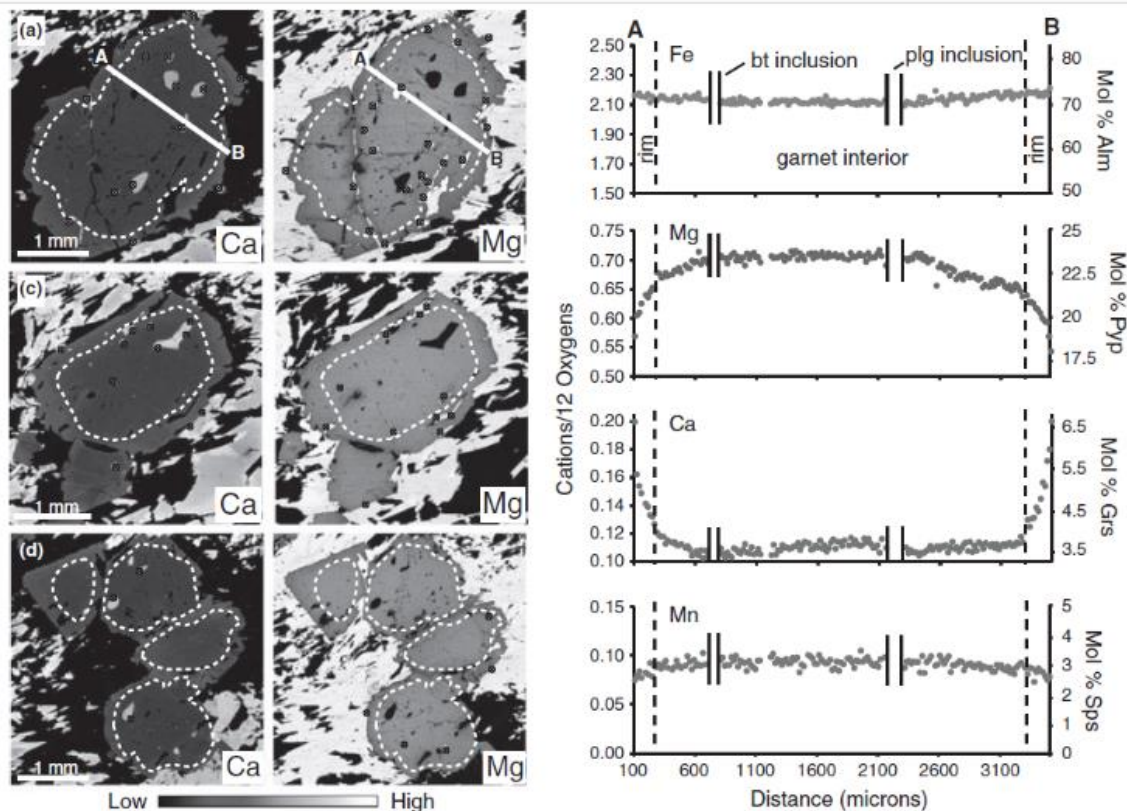


Figure 2: An x-ray map showing the geochemistry of a garnet porphyroblast from the Northern Appalachians. This zoning was the result of partial breakdown of garnet rims caused by contact metamorphism.

A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

75 Biotite is distributed throughout the matrix as anhedral crystals and also as inclusions within
76 garnet poikiloblasts. Compared to the matrix plagioclase, which shows distinct calcium zoning,
77 there is no evidence of iron or magnesium zoning of biotite regardless of its location. Spinel is
78 only present within the matrix and not as inclusions within the core of garnet crystals (Dorfler,
79 Tracy, & Caddick, 2014).

80 Multiple periods of metamorphism are also suggested by the lack of zoning in inclusions and
81 absence of spinel in some samples (Dorfler, Tracy, & Caddick, 2014). This can also be used to
82 extrapolate how some metamorphism only affected minerals in the matrix (Dorfler, Tracy, &
83 Caddick, 2014). The presence of spinel is also an indicator of high-temperature metamorphism.
84

85 Central Appalachians and the Acadian Orogeny:

86 The central Appalachians contain the Taconic and Catskill Mountains, as well as the northernmost
87 part of the Blue Ridge (see Figure 1) (Marshall University, 2010). Beginning in the Devonian, the
88 Acadian orogeny formed part of the northern and central Appalachians, but petrologic and dating
89 evidence has shown limited effects in the southern chains (Murphy, n.d.; Marshall University,
90 2010; Rast, 1989). The majority of the central Appalachian chain was made by mid-carboniferous
91 greenschist to lower-amphibolite facies metamorphism (Murphy, n.d.; Rast, 1989).

92 The northern Blue Ridge basement is composed of granitic gneisses and weakly to strongly
93 foliated granitoids. The gneisses are typically
94 older than the surrounding granitoids, and the
95 crystallization of most rocks indicates a
96 Granulite facies (Tollo, Aleinikoff, Borduas,
97 Hackley, & Fanning, 2004).

98 In order to determine the geographical extent
99 to which the northernmost part of the Blue
100 Ridge was affected by the Acadian Orogeny,
101 Tollo, Aleinikoff, Borduas, Hackley, & Fanning
102 (2004) examined samples from two separate
103 massifs on opposite sides of a fault that runs
104 parallel to the Appalachian chain. Rocks located
105 on either side of the fault are coarse-grained
106 with porphyroblasts. The eastern side of the
107 fault shows a slightly lower grade of
108 metamorphism than the western side. Biotite is
109 the dominant mineral in eastern gneisses, and it
110 shows varying degrees of recrystallization.
111 Orthopyroxene and garnet are the major
112 minerals in the western samples, which show a
113 stronger gneissose texture when compared to
114 their eastern companions. The oldest rocks,
115 located on the eastern side of the fault, are
116 strongly foliated granitic/monzonitic gneisses
117 (Tollo, Aleinikoff, Borduas, Hackley, &
118 Fanning, 2004).

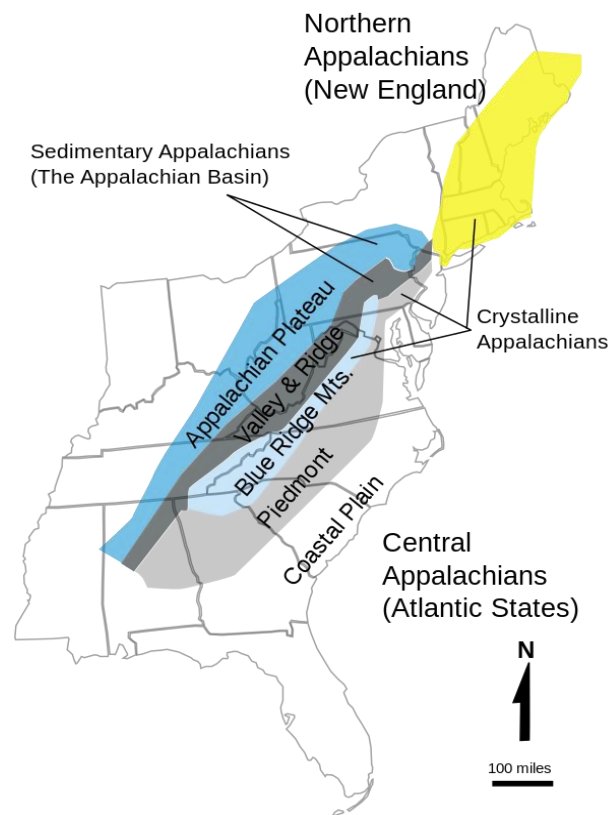


Figure 3: A map showing the Blue Ridge Mountains and Piedmont province in the Northern and Central Appalachians.

A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

119 Two hypotheses were proposed to explain the varying degrees of metamorphism:
120 1. The basements of these two massifs are two different responses to amphibolitic
121 metamorphism
122 2. The two sides of the fault experienced different degrees of metamorphism
123 Additional to having experienced a higher grade of metamorphism, the rocks on the western
124 side of the fault are slightly younger and possess low amounts of dark minerals when compared to
125 the eastern side. This makes the second hypothesis more likely to explain the variations (Tollo,
126 Aleinikoff, Borduas, Hackley, & Fanning, 2004).

127 Southern Appalachians and the Alleghenian Orogeny:

128 The third orogeny, which formed most of the southern Appalachians is the Alleghenian orogeny
129 of the Late Carboniferous and Permian, also known as the Appalachian Orogeny (Murphy, n.d.;
130 Marshal University, 2010; Hatcher, 2002). The entirety of the Appalachian chain shows the
131 regional metamorphism of this event as it was
132 caused by the rotational collision of the African
133 continent with the eastern coast of North America
134 (Murphy, n.d.; Rast, 1989; Hatcher, 2002). Figures
135 3 and 4 were adapted from PerryGeo (2006) and
136 Merrigum (n.d.), and they show that North America
137 and Africa were part of larger continents called
138 Laurentia and Gondwana during this orogeny
139 (Hatcher, 2002). This tectonic event can be
140 classified as what is known as “Zipper Tectonics,” where the
141 Northern edge of Africa collided with Eastern North
142 America (Murphy, n.d.; Rast, 1989; Hatcher, 2002).
143 In the process of this collision, the Theic Ocean
144 between these two continents was closed from North
145 to South in a zipper-like pattern, causing subduction-
146 setting metamorphism (Hatcher, 2002). This model
147 explains why similar metamorphism is seen
148 throughout the Appalachian Mountains but with
149 different age ranges (Hatcher, 2002). Several I-Type
150 and S-Type intrusions were also created in this
151 collision and caused additional contact
152 metamorphism (Rast, 1989; Hatcher, 2002). While
153 there was a large amount of time between the first
154 two orogenies that created the northern and central
155 Appalachians, the Alleghanian orogeny almost
156 immediately followed the preceding Acadian
157 orogeny. (Rast, 1989). This causes difficulty when
158 attempting to differentiate between the regional
159 metamorphism caused by each event, especially near
160



Figure 4: A map of Laurentia with an outline of modern North America. The red zone represents the Appalachian Mountains.



Figure 5: A map of Gondwana with inner traces of the continents it was composed of. The red patch is where Africa contacted North America in the Appalachian Orogeny.

A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

161 the boundary between the central and southern sectors.

162 The southern Appalachians encompass the Blue Ridge and Great Smokey Mountains (see
163 Figure 1) (“USGS: Geology Of The Southern Appalachian Mountains,” n.d.). The Blue Ridge and
164 Piedmont province on the eastern side of the Appalachians were the main areas affected by this
165 orogeny (see Figure 3). The Piedmont province, in contrast to the greenschist facies of the Blue
166 Ridge, shows higher-degree amphibolite facies (Williams, 2011; Hatcher, 2002). Hatcher (2002)
167 found the upper-amphibolite facies to have reached sillimanite conditions. There are also
168 differences in basement composition throughout the Blue Ridge, with the north being comprised
169 of gneisses and the south of charnockites (Rast, 1989).

170 Chunky Gal Mountain is a highly-publicized massif within the Blue Ridge whose basement
171 experienced amphibolite facies metamorphism (Marshall University, 2010). Williams (2011)
172 described the minerals and textures from opposite sides of the Chunky Gal Mountain fault to
173 characterize the metamorphism of the central Blue Ridge. The study discovered that the more
174 southerly located samples displayed lower degrees of metamorphism (amphibolitic) than those
175 more northern (granulite) (Williams, 2011). This may infer that the Alleghanian orogeny had more
176 limited effects near the southernmost part of the Appalachian chain, as the grade of metamorphism
177 decreases towards the furthest edges of the Southern section (Hatcher, 2002; Williams, 2011).

178 Garnet and plagioclase crystals contained the most geochemical evidence for peak pressure and
179 temperature values, as they would preserve the most evidence of changing conditions. Iron and
180 magnesium oxide data in garnet and biotite were used for thermobarometry, while aluminum and
181 silicate oxide data from garnet and plagioclase was used to calculate barometry values. These
182 values were determined to range from 5.5-8kbar depending on the area examined, showing the
183 differences in metamorphic grade with southern movement along the Blue Ridge. (Williams,
184 2011).

185 Garnet crystals in the examined thin sections have cores abundant in inclusions, but rims only
186 contain trace amounts of sillimanite. Garnets examined showed zoning, poikiloblastic textures,
187 and fractionation, which were evidence of three separate growth stages followed by a period of
188 cooling (Williams, 2011; Hatcher, 2002):

- 189 1. Growth of garnet porphyroblasts that encompasses some of the surrounding mineral
190 grains
- 191 2. Outward expansion of the garnet during a period where there was no remaining material
192 to form inclusions
- 193 3. A period where sillimanite inclusions were incorporated into the garnet rim and tail
- 194 4. A rapid cooling period that caused the garnet crystals to fracture

195 Trace amounts of sillimanite inclusions within the rims and tails of the poikiloblasts differ from
196 those in the matrix. A more fibrous shape is seen in the inclusions while matrix sillimanites have
197 a more prismatic appearance. Since sillimanite inclusions are parallel to the crystals in the matrix,
198 it is likely that they were formed in the same deformation event (Williams, 2011).

199 A green mica, along with fine crystals of potassium feldspar, was identified within the fractures
200 of garnet poikiloblasts. As it is within the fractures and not the garnet grains themselves, it can be
201 inferred that they formed after the porphyroblasts and during the final period of metamorphism in
202 which the grains were cooling. The green mica, although not named in the study, is likely chlorite
203 based on its birefringence and green color in plane polarised light (PPL) (Williams, 2011).

A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

204 Potassium feldspar is also contained within the fractures and rim of garnet poikiloblasts. It
205 lines the rim of the poikiloblast with a myrmekitic texture, potentially indicating an interaction
206 with hydrothermal fluids at some point during the metamorphism (Williams, 2011).

207 **Summary and Conclusions:**

208 The Taconic orogeny was the first mountain-building event to occur, forming the northern
209 Appalachians (Murphy, n.d.; Marshall University, 2010; Hatcher, 2002). Because of this it was
210 also repeatedly impacted by mountain-building events, making it difficult to distinguish between
211 the metamorphic results of each orogeny in some areas of the northern ranges. In this sector,
212 regional metamorphism was later overlain by contact metamorphism. This causes a discrepancy
213 in the mineralogy of samples that experienced one or both types of metamorphism. Multiple
214 periods of metamorphism can also be seen in alignment and recrystallization of matrix minerals
215 (Dorfler, Tracy, & Caddick, 2014). The main evidence of this is seen in the inclusions of garnet
216 poikiloblasts and presence of matrix spinel and sillimanite.

217 The Acadian orogeny followed the Taconian orogeny and created the central and part of the
218 northern Appalachians (Murphy, n.d.; Marshall University, 2010; Rast, 1989). The more eastern
219 central Appalachians were determined to be older but experienced less metamorphism, which may
220 explain why biotite is the dominant mineral in the area. The western central Appalachians,
221 although younger, are of a higher grade and contain high amounts of orthopyroxene and garnet.
222 Differences between each side of the fault is most likely the result of experiencing different degrees
223 of metamorphism followed by thrusting.

224 The Alleghanian orogeny quickly followed the Acadian event and affected the entirety of
225 Appalachians (Rast, 1989; Hatcher, 2002). It caused extensive regional metamorphism and minor
226 amounts of contact metamorphism from the mafic intrusions it created in some areas (Hatcher,
227 2002). It is also the only sector to show a potential for interaction with hydrothermal fluids.

228 The largest discrepancies in mineralogy and degree of metamorphism are between the
229 northern and southern Appalachians. Cooling of northern garnet grains was intermittent between
230 two types of metamorphism, but cooling in southern samples allowed for extensive mineral growth
231 in fractures that did not occur in northern counterparts (Murphy, n.d.; Aleinikoff, Borduas,
232 Hackley, & Fanning, 2004). Although similar in a lack of inclusions, the outer rims of garnet
233 poikiloblasts in the north and south were not made by the same processes. Northern garnets have
234 a rim that was made through the recrystallization of the outer crystal without any increase in size,
235 while the rim of southern garnets represents growth with a lack of surrounding material for
236 inclusions.

237 Further research could be improved with increased usage of an electron probe to give more
238 accurate results than a scanning electron microscope, which focusses on a broader area of the thin
239 section. This would better provide peak metamorphic temperature and pressures (Williams, 2011).
240 Future studies may benefit from a research focus that distinguishes and defines the geographic
241 boundaries of each orogeny, which would better explain the varying metamorphic facies
242 throughout similar massifs.

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A REVIEW OF THE METAMORPHIC PETROLOGY OF THE APPALACHIAN MOUNTAINS

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